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## 1960 through 1972, a Time of Discoveries in Large-Scale Tropical Meteorology

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**Abstract.** The Australian Bureau of Meteorology ([bom.gov.au](http://bom.gov.au)) states that “The Madden-Julian Oscillation (MJO) is a major fluctuation in tropical weather on weekly to monthly timescale. The MJO can be characterized as an eastward moving ‘pulse’ of cloud and rainfall near the equator that typically recurs every 30 to 60 days.” Early descriptions of the MJO were contained in two papers by Madden and Julian (1971; 1972). This paper relates the story of developments in tropical meteorology in the 1960s that led to those two papers. The decade saw the first unambiguous identification of large-scale, theoretically predicted, tropical waves. Spectral analysis was used effectively by researchers to link observations with the theoretically expected features of these waves. At the same time, longer time series of observations, faster computers, and an algorithm designed to speed up Fourier transforms, vital for spectral analysis, all became available. These developments set the stage for the oscillation to be recognized.



1   **1) Introduction**

2

3       Clarence Palmer, in his treatise on Tropical Meteorology, stated  
4 that whenever we get more data from the tropics “...the results  
5 usually astonish us” (Palmer, 1952). That certainly applied to the  
6 decade of the 1960s. Tropical data were becoming more accessible  
7 in digitized form suitable for treatment by computers then also  
8 available to meteorologists. Proceeding logically from Palmer’s  
9 observation, the stage was set for us to be astonished and  
10 astonished we were.

11

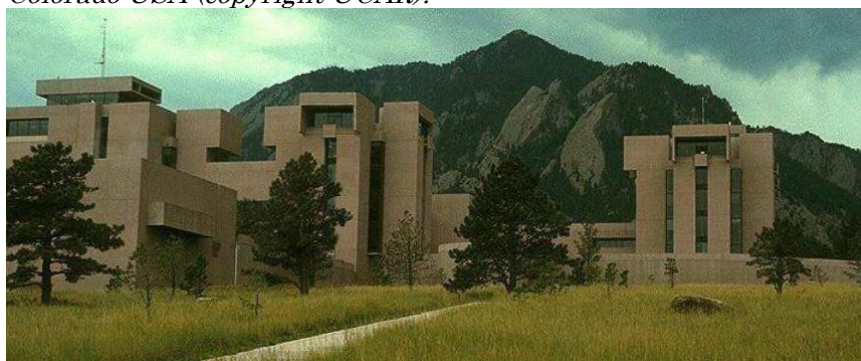
12       The decade opened with the discovery of the Quasi-Biennial  
13 Oscillation (QBO) in the equatorial stratosphere. It is a most  
14 amazing phenomenon. The QBO is remarkable in its approximate  
15 26-month period, its regular downward propagation, and in its  
16 large amplitude. By the mid-1960s efforts to explain the QBO led  
17 to two of the earliest unambiguous identifications of large-scale  
18 atmospheric waves predicted by theory. Simultaneously, a theory  
19 tailored just for tropical regions was published. Spectral analysis, a  
20 particularly powerful analysis tool for the tropics that sometimes  
21 requires relating events at stations 1000s of kilometers apart, was  
22 beginning to be used effectively by researchers. Also, in 1965 a  
23 fast Fourier transform algorithm suitable for coding was published  
24 that made spectral calculations orders of magnitude faster than  
25 traditional ones.

26

27       Toward the end of the decade in early 1967 a large-scale field  
28 program in the Equatorial Central Pacific was organized and  
29 carried out by the National Center for Atmospheric Research.

30

31       Figure 1. *National Center for Atmospheric Research, Boulder,*  
32 *Colorado USA (copyright UCAR).*



33

34



1 Work on data from this field program, the Line Islands Experiment  
2 (LIE), introduced me, a rookie researcher at NCAR, to the  
3 developments described above. In the fall of 1967, during my first  
4 days at NCAR, I sat in the temporarily vacant office of an NCAR  
5 scientist who was spending a sabbatical year at the University of  
6 Chicago from where, coincidentally, I had just left fresh with a  
7 Masters Degree. That scientist was Paul Julian. We were both  
8 now members of NCAR's Synoptic Meteorology Group and later the  
9 Empirical Studies Group (Fig. 2).

10

11 Figure 2. *Empirical Studies Group: left to right, Paul Julian,*  
12 *Roland Madden, Dennis Shea, Chester Newton (Group Head), and*  
13 *Harry van Loon about 1980 (copyright UCAR).*

14



15

16

17 Julian was a former student of Hans Panofsky who pioneered  
18 the use of spectral analysis in meteorology (Panofsky, 1955). As a  
19 result, Julian was well versed in the technique. He published a  
20 review of it and showed how spectral analysis could be applied to  
21 geophysical data (Julian, 1967). He had also done a spectral  
22 analysis of 21 years of zonal index data (a measure of the strength  
23 of the westerlies) to test quantitatively the notion of an index cycle  
24 of three to eight weeks (Julian, 1966) that had been qualitatively  
25 accepted as real (e.g. Petterssen, 1956). Results suggested that the  
26 zonal index could be adequately modeled by a first order



1 autoregressive process with little evidence of a preferred three to  
2 eight week variation.

3  
4 Upon Julian's return to NCAR in 1968 he guided me in my  
5 efforts to apply spectral analysis to the LIE winds. By late 1970,  
6 our collaboration had expanded and eventually led to a description  
7 of what is now referred to as the MJO, or Madden and Julian  
8 Oscillation after two papers in the early 1970s (Madden and  
9 Julian:1971; 1972; hereafter MJ71 and MJ72). What follows is the  
10 story, as I remember, of our collaboration that led to those two  
11 papers and of the background outlined above that influenced us.

## 12 13 **2) Background**

### 14 15 **2-1) Discovery of the Quasi-Biennial Oscillation**

16  
17 Graystone presented a time-height diagram of zonal, or u-wind,  
18 in the stratosphere over Christmas Island, 2N and 158W  
19 (Graystone, 1959). He had only two years of data and could not  
20 recognize the amazing QBO which he was sampling. He only  
21 remarked that there was an absence of annual cycle and a presence  
22 of large vertical shears. Ebdon (1961) and Reed et.al. (1961)  
23 extended the record beyond two years, and they were able to  
24 identify east and west wind regimes propagating downward in the  
25 stratosphere that varied by as much as 40 m/s. The average time  
26 scale of the wind shifts was 26 months. I am sure that Palmer, who  
27 left us in 1973, was astonished along with most meteorologists by  
28 the amazing behavior of the QBO. The QBO increased interest in  
29 tropical meteorology. Although the QBO did not affect our later  
30 work directly, some subsequent research aimed at explaining its  
31 behavior did.

### 32 33 **2-2) Theory of Waves in the Equatorial Atmosphere**

34  
35 Matsuno (1966) published a theoretical paper specifically tailored  
36 to the equatorial region. He showed how his approximate  
37 equations are an asymptotic case of the Laplace Tidal Equations  
38 (LTE) that address the behavior of a thin fluid on the full, rotating,  
39 spherical earth. The wave solutions of the LTEs fall into two  
40 classes (Hough, 1898): The waves of the First Class are eastward  
41 and westward traveling gravity waves; waves of the Second Class  
42 are westward propagating, and called Rossby or normal mode  
43 Rossby-Haurwitz waves (Rossby et.al., 1939; Haurwitz, 1940a ;  
44 1940b). Similarly, Matsuno's equations yield two classes of waves,



1 inertia gravity waves and Rossby waves, which are equivalent to  
2 approximate forms of waves of the First and Second Class.

3  
4 Two waves, or modes, described by Matsuno are of special  
5 interest. One mode behaves like a gravity wave, or wave of the  
6 First Class, for waves long relative to the fluid depth and like a  
7 Rossby wave, or wave of the Second Class, for shorter waves (see  
8 Longuet Higgins, 1968: Fig. 5 for zonal wavenumber 4). This mixed  
9 Rossby gravity wave (MRGW) is reflected in variations of the  
10 meridional, or v-wind.

11  
12 The second important mode is a special type of wave of the First  
13 Class, the atmospheric Kelvin wave. Unlike the MRGW, the  
14 equatorial Kelvin wave is confined to variations in the u-wind. It  
15 gets its name from work by Lord Kelvin (Thomson, 1879) who  
16 studied waves that propagate parallel to sides of a canal. The  
17 interesting aspect of atmospheric, equatorial Kelvin waves is the  
18 fact that the change in sign of the Coriolis force at the equator acts  
19 dynamically as the canal side (or a coast line).

20  
21 In his 1966 paper, Matsuno posed the question of whether the  
22 waves he described exist in actual atmospheric conditions. The  
23 answer came back quickly in research aimed at explaining the  
24 QBO.

### 25 26 **2-3) Discovery of MRGWs and Kelvin Waves in Observations**

27  
28 Yanai and Maruyama (1966), searching for evidence of eddy  
29 disturbances that might converge enough momentum to drive the  
30 QBO, discovered alternating downward propagating north and  
31 south winds in the stratosphere over the central tropical Pacific  
32 during northern spring and early summer of 1958. The average  
33 period of the oscillating v-winds was around five days. These  
34 varying winds were shown to behave similarly to those of MRGW  
35 (Maruyama, 1967).

36  
37 The discovery of Kelvin waves soon followed (Wallace and  
38 Kousky, 1968). Like Yanai and Maruyama (1966), Wallace and  
39 Kousky were initially motivated by problems related to the  
40 momentum budget of the QBO. They found fluctuations in the u-  
41 wind in the stratosphere at stations in the Pacific and Caribbean  
42 with an average period of 15 days. Wallace and Kousky showed  
43 that the structure and behavior of the oscillations were consistent  
44 with those of the theoretically predicted equatorial Kelvin wave.

45 .



1 It should be said that the discovery of MRGW and Kelvin waves  
2 were among the first unambiguous identifications of large-scale  
3 atmospheric waves predicted by theory. With the exception of the  
4 findings of Kubuto and Iida (1954) that showed the presence of  
5 normal mode Rossby-Haurwitz waves, in the mid-60s there was  
6 little observational evidence of theoretically predicted large-scale  
7 waves. The important papers of Eliason and Machenhauer (1965;  
8 1969) identifying normal mode Rossby-Haurwitz waves were either  
9 just being disseminated or still on the drawing boards, as was Ray  
10 Deland's work (e.g. Deland, 1965).

11

#### 12 **2-4) Use of Spectral Analysis in Studying Tropical Data**

13

14 Besides showing the similarity between Yanai-Maruyama waves  
15 and Matsuno's MRGWs, Maruyama (1967) used spectral analysis in  
16 the diagnosis of the data. It showed spectral peaks, or extra  
17 variance, in the 4-5 day period range quantifying the subjectively  
18 estimated period in Yanai and Maruyama (1966).

19

20 Wallace and Kousky (1968) used spectral analysis to identify  
21 aspects of the theoretical Kelvin wave in the upper air data that  
22 they were examining. For example, cross-spectra between zonal  
23 wind and temperature quantified a quadrature relationship  
24 predicted by theory and underscored the power of spectral analysis  
25 when diagnosing wave-like behavior.

26

27 Yanai and colleagues at the University Tokyo and Wallace and  
28 colleagues at University of Washington then expanded their use of  
29 spectral analysis to further diagnose tropical wave motions. Yanai  
30 et.al. (1968) computed spectra and cross-spectra of the v-wind  
31 between 17 Pacific stations at 34 levels from the surface to the  
32 lower stratosphere. Data were from the period April to July of 1962.  
33 The analyses allowed them to estimate vertical and horizontal  
34 structures of 4-5 day period disturbances. Among other things, the  
35 found that in the lower troposphere, the v-wind spectra had  
36 spectral peak near 4-days and phase angles that suggested an  
37 eastward slope with height.

38

39 Wallace and Chang (1969) studied data from the July to  
40 December 1963 period. They concentrated on the troposphere  
41 below 500hPa. They saw little evidence of vertical propagation in  
42 the 4-5 day v-wind variations in contrast to that in the 1962 data  
43 examined by Yanai et.al. (1968). Wallace and Chang looked at  
44 three additional six month periods during the two years 1963  
45 through 1964 at Truk Island (7N, 152E, now Chuuk). They



1 determined that besides the vertical structure changing with time,  
2 the 4-5 day spectral peak in the v-wind itself varied with time as  
3 well.

4  
5 Wallace and Chang also detected the 5-day normal mode Rossby-  
6 Haurwitz wave that had recently been identified by Eliassen and  
7 Machenhaur (1965; 1969) and Deland (1965). Most important for  
8 the MJO story is that they reported on a low frequency oscillation  
9 "...which could not be adequately resolved with the limited period  
10 of record" (Wallace and Chang, 1969). Like Graystone's limited  
11 look at the QBO, Wallace and Chang's low frequency oscillation  
12 may have been the MJO awaiting a longer record to be recognized.

13  
14 Spectral analyses by Yanai, Wallace, and colleagues had a major  
15 influence on our later studies. Much of the work described above  
16 was summarized by Wallace (1969) along with examples of the use  
17 of spectral analysis. Innovative uses of spectral analyses of these  
18 tropical data were further summarized by Wallace (1971) and by  
19 Julian (1971).

## 20 21 **2-5) The Fast Fourier Transform**

22  
23 Spectral analysis and cross-spectral analysis involves Fourier  
24 transforming the time series data directly (direct method), or  
25 Fourier transforming the auto-covariance function determined from  
26 the time series data (indirect method). In either case the  
27 traditional Fourier transform requires considerable multiplications.  
28 In 1965, Cooley and Tukey (1965) published an algorithm suitable  
29 for computer calculations of a fast Fourier transform (FFT). The  
30 FFT sped up computations enormously. For a time series N values  
31 long the traditional transform required  $N^2$  multiplications.  
32 Depending on how factorable N is, the FFT required about  
33  $N \cdot \log_2(N)$  or a speed up factor of  $N/\log_2(N)$  (Cooley and Tukey,  
34 1965; Cooley, 1987). For  $N=1000$  the speed up factor is 100. So the  
35 door was opened for much faster transforms.

36  
37 It is interesting to note that Cooley reports that his first  
38 interaction with Tukey was at Princeton's Institute for Advanced  
39 Study Program in 1953 where Tukey was a consultant and he,  
40 Cooley, was a programmer in Von Neuman's Numerical Weather  
41 Prediction Group. Cooley programmed a spectral analysis routine  
42 for Tukey (Cooley, 1987). Ten years later the two would team up to  
43 forever alter the way we compute a Fourier transform.

## 44 45 **3) Work Leading to the Identification of the MJO**



1  
2 **3-1) The Line Islands Experiment Upper Air Data and the**  
3 **Honolulu Tropical Meteorology Conference of 1970**  
4

5 The above happenings laid the groundwork for Julian and my  
6 work that led to a description of the MJO. The role that spectral  
7 analysis played in the discoveries of MRGWs and Kelvin waves,  
8 and its demonstrated value describing tropical, tropospheric  
9 disturbances suggested that we should apply it to the LIE data..  
10

11 Backing up a little, in early 1967 I was anticipating completing  
12 the requirements for a Masters Degree at the University of  
13 Chicago. I had spent the prior two years studying under Professor  
14 Tetsuya Fujita and had learned a lot about Satellite Meteorology.  
15 Given a satellite's attitude in space and its subpoint, I had learned  
16 to "grid" or add latitude/longitude lines to any picture. This skill  
17 put me in a good position to qualify for a job opening at the newly  
18 started NCAR. NCAR was in the process of carrying out the LIE  
19 under the direction of Chief Scientist Ed Zipser. The LIE was  
20 motivated in part to provide ground truth for ATS-1, the first  
21 equatorial geosynchronous meteorological satellite. It was  
22 launched in December 1966 (Zipser, 1970). When I arrived in  
23 Boulder in September of 1967, my first assignment was to grid  
24 pictures taken by ATS-1 earlier that year during the LIE.  
25

26 As the picture gridding neared completion, probably in late  
27 October 1967 a large shipping crate appeared in the hall outside my  
28 office. I learned that it was full of punch cards containing  
29 thermodynamic, azimuth, and elevation information of more than  
30 800 rawinsondes recorded during the LIE. It was then my  
31 responsibility to turn these raw data into wind speed and direction,  
32 temperatures, and moisture variables. Fortunately, NCAR  
33 scientist Ed Danielson and summer student Bob Gall, who years  
34 later would become a Division Director at NCAR, had written a  
35 computer program to do just that for the LIE data. I teamed with  
36 Dennis Joseph, a data expert and member of NCAR's Computing  
37 Facility's Data Support Section, to finish the job. The  
38 thermodynamic and wind data were published in February 1971  
39 (Madden et.al., 1971).  
40

41 In the meantime, we had the opportunity to look at the data. I  
42 am not certain, but I think it is likely that when I arrived at NCAR  
43 in 1967, I was unaware of the important discoveries summarized  
44 above. Probably in early 1968, one of my new NCAR colleagues  
45 drew my attention to the Yanai and Maruyama (1966) paper





1 because of its relevance to LIE stratospheric data. .Certainly by  
2 early 1969, I had learned of the innovative ways that Yanai,  
3 Wallace and colleagues were using spectral analysis in their work.  
4 It was natural to do similar analyses for the LIE period.

5  
6 By mid-1968, I had been moved to a more permanent office and  
7 Julian had returned from Chicago. In the months to follow, with  
8 his help, I computed spectra of LIE upper air meridional winds and  
9 was preparing a paper to be delivered at the Honolulu Tropical  
10 Meteorology Meeting planned for June of 1970. Figure 3 is a  
11 photograph of meeting participants (Ramage, 1970). Many who are  
12 mentioned in the text are circled. My paper was entitled “Wave  
13 Disturbances over the Equatorial Pacific during the Line Islands  
14 Experiment” (Madden, 1970). Results did not show the 4-5 day  
15 spectral peaks in the lower troposphere that were present during  
16 April-July of 1962 reported by Yanai et.al. (1968) further  
17 confirming the variability of the tropospheric spectra.

18  
19 Figure 3. *Participants in the Honolulu Conference, 1970. (photo*  
20 *credit American Meteorological Society/World Meteorological*  
21 *Organization, 1970). Key to all in the photo can be seen in Ramage*  
22 *(1970) or Bulletin of the American Meteorological Society*  
23 *Supplemental Materials (Madden, 2019).*

24



Participants in the Honolulu Conference (from Proc. of the Sym. on Tropical Meteorology, Honolulu, HI, 1970).

25  
26



1 In the discussion that followed my talk, Gary Atkinson from the  
2 U. S. Air Force made the observation that the nascent spectral  
3 studies of the tropics had been based on relatively short time series  
4 (the LIE time series were only 47 days long). He suggested that  
5 there now were long time series available from Pacific stations and  
6 it would be good to compute spectra based on them to better assess  
7 the "...variability and/or stability..." of results. We knew that  
8 Atkinson was right and that Julian and I were in the perfect  
9 position to look at longer time series. The NCAR Data Support  
10 Section headed by Roy Jenne had begun to collect some of these  
11 long series. Julian had a FFT code based on Cooley and Tukey's  
12 algorithm, and we had access to a Control Data Corporation 6600.  
13 The CDC 6600 had a clock speed of 10mHz and a memory of 64kb,  
14 which though orders of magnitude slower and smaller than a  
15 modern cell phone, made it the most powerful computer available in  
16 the 1960s for meteorological research. Upon our return from  
17 Honolulu I turned my attention to investigating the longest time  
18 series available from the equatorial region.

19  
20 Figure 4. *CDC 6600 (copyright UCAR).*



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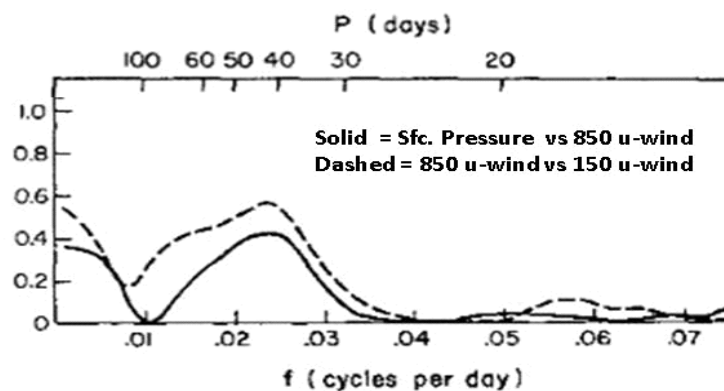
### 3-2) Studying Longer Time Series

#### 3-2-1) MJ71



1  
2 Our motivation was to examine time variation in the spectra of  
3 tropical observations in the 4-5 day period range. The longest  
4 record available for this purpose was rawinsonde data from 3584  
5 days measured at Canton Island (3S, 172W, now Kanton). With the  
6 long record we could resolve lower frequency variations that had  
7 not yet been investigated. Almost immediately our attention  
8 shifted from documenting time variations in 4-5 day disturbances to  
9 investigating variations in the 40-50 day range because of results  
10 typified by Fig. 5. Coherence squared shown in Fig. 5 is similar to  
11 correlation as a function of frequency. It shows a broad maximum  
12 with largest values in the 40-50 day range. Corresponding phase  
13 angles (not shown) indicated that surface pressures and 850 hPa u-  
14 winds were in phase and 850 and 150 hPa u-winds were out-of-  
15 phase.

16  
17 We had no a priori reason to expect this result so usual  
18 statistical tests were not appropriate. Julian discussed prior versus  
19 posterior statistical tests in the paper we prepared (MJ71), and  
20 demonstrated that it would be rare to have such high coherence  
21 values if the time series were not related. He expanded on this  
22 argument in his paper published in 1971 (Julian, 1971).



23  
24 Figure 5. *Coherence squared between 850 and 150 u-winds*  
25 *(dashed) and between surface pressure and the 850 u-winds (solid)*  
26 *based on Kanton pressure and upper air data . The 0.1% prior*  
27 *confidence level assuming a null of no Coherence is 0.25. (adapted*  
28 *from MJ71)*  
29

30 Considering phase angles between participating variables we  
31 concluded the evidence pointed to a large circulation cell orientated  
32 in the equatorial plane with a node where the u-wind switches



1 direction in the 600-500 hPa region. Neither the spectra nor cross-  
2 spectra involving the v-wind showed maxima in the 40-50 day  
3 range so we concluded that it was not involved. We will see that  
4 data from more stations confirmed that the oscillation was indeed  
5 the manifestation of large circulation cells, but, because none of our  
6 spectra differentiated between seasons, our conclusion about v-  
7 wind was wrong.

8  
9 The paper describing the results from Kanton was submitted to  
10 the Journal of the Atmospheric Sciences on 21 December 1970. A  
11 letter dated 12 March 1971 from editor S. I. Rasool stated that  
12 reviews, plural, were in but relatively minor comments from only  
13 one review were attached. We addressed the reviewer's comments  
14 and the paper was published in July 1971 (MJ71).

### 15 16 **3-2-2) MJ72**

17  
18 Working to get a full picture of the phenomenon that we were  
19 seeing at Kanton, we assembled station pressure data from 25  
20 tropical stations (16 were within  $15^{\circ}$  of the equator) and upper air  
21 data from six stations all located within  $15^{\circ}$  of the equator and  
22 spaced around its full circumference. Cross-spectra of the  
23 pressure data revealed that the 40-50 day disturbance propagated  
24 eastward and spread poleward but was strongest within  $10^{\circ}$  of the  
25 equator. Based on cross-spectra of upper air data from the six  
26 stations, we made a figure that summarized the status at all  
27 stations and levels that were coherent with Kanton pressure when  
28 Kanton pressure was a relative maximum (Fig. 6 of MJ72). That  
29 helped us to envision the zonally orientated circulation cells and  
30 their eastward movement.

31  
32 To supplement the spectral evidence with a synoptic picture of  
33 the disturbance based directly on time-series data, we turned to  
34 data from the International Geophysical Year, 1957-1958. (IGY).  
35 Willy Rudloff had presented a paper at the Honolulu meeting titled  
36 "Measurable Seasonal Variations in the Total Mass of the  
37 Atmosphere" (Rudloff, 1970) which was based on gridpoint pressure  
38 data digitized from IGY World Weather Maps prepared by his  
39 office, the Seewetteramt in Hamburg. I wrote to him on 12 March  
40 1971 about our interest in the tropical zone and he kindly sent us  
41 the grid point sea level pressure data on computer cards which was  
42 a standard way of storing and transferring data. We also tabulated  
43 and prepared punched cards containing IGY upper air data from  
44 printouts available in the NCAR library. In 1970, all of our  
45 programs and much of our data were contained on punch cards.



1 Typical of the time, Fig.6 shows two NCAR programmers  
2 submitting their card decks to be read into the CDC 6600.

3

4

5 Figure 6. *Two programmers (left) submitting their card decks to the*  
6 *operator for reading into the CDC 6600 about 1970 (copyright*  
7 *UCAR).*

8



9

10

11 We computed a composite wave by first selecting dates during  
12 the IGY period when 45-day band-pass filtered Kanton pressure  
13 was at a relative minimum, and separately at a maximum and at  
14 six more intermediate times. The IGY data were then averaged for  
15 each of the set of eight dates separately. A picture emerged that  
16 was consistent with the spectral results. The sea level pressure  
17 perturbations moved eastward as did those of the zonal wind.  
18 There was a wave on the tropopause and some evidence of water  
19 vapor mixing ratio variations consistent with eastward moving  
20 deep convection. A more detailed discussion of the first time we  
21 saw eastward propagation in the IGY pressures is contained in  
22 Hand (2015).

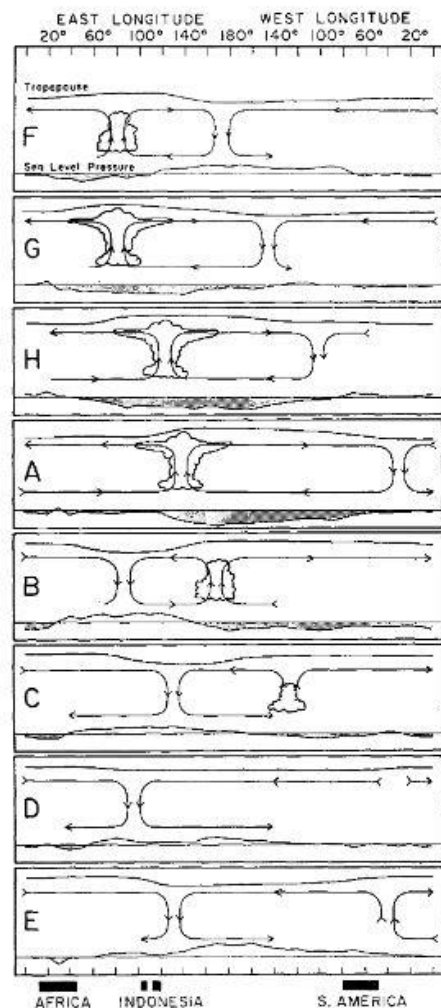
23

24 This spectral and synoptic evidence led to a description of the  
25 40-50 day oscillation that is contained in Fig. 7. Phase E



1 corresponds to the time when station pressure is a relative  
2 maximum at Kanton and Phase A when it is a minimum. We had  
3 no precipitation or cloud data at the time but included an indication  
4 of varying convection because of the low level convergence in the u-  
5 wind, mixing ratio changes, and the changing height of the  
6 tropopause.

7  
8 *Figure 7. Schematic depiction of the time and space variations of*  
9 *the oscillation. Phase “A” (“E”) is the time of lowest (“highest”)*  
10 *pressure at Kanton. Other phases are intermediate times and for a*  
11 *48-day period are approximately six days apart. Anomaly*  
12 *pressures are indicated at the bottom of each panel with negative*  
13 *anomalies shaded. Circulation cells are based on u-wind*  
14 *variations. Regions of enhanced convection are proposed based on*  
15 *u-wind convergence/divergence, tropopause height differences (top*  
16 *line), and mixing ration changes.*  
17



1  
2 We submitted our paper describing the above results on 6 April  
3 1972. Editor Rasool, in a letter dated 8 May 1972, stated that "...  
4 your paper has been found acceptable for publication ..." This time  
5 comments from two reviewers were included. Reviewer 1 accepted  
6 the paper on the condition that results from Gan Island (0.7S,  
7 73.2E) are included. The reviewer stated that "... Gan Island is  
8 strongly affected by the Asian Monsoon which seems to possess a  
9 30-40 day period", a qualitative assessment which was  
10 quantitatively documented by Yasunari (1979; 1980).

11  
12 Fortunately, during the review process we had begun to examine  
13 spectra and cross-spectra for Gan data, and results were easily  
14 added to Figs. 1 and 4 of MJ72. Reviewer 2 gave us eight



1 constructive comments which required only small changes. The  
2 paper was published in September 1972. We pointed out that the  
3 oscillation was a broad-band one, but called it the “40-50 Day  
4 Oscillation” because spectral maxima of the various variables most  
5 often fell in that range. The “MJO” reference began being used  
6 more frequently after it appeared in the title of two papers  
7 (Swinbank et. al., 1988; Lau et. al., 1988).

8  
9 Table 1 shows the sequence of submission and publication dates  
10 for some of the relevant papers. Recently, Li et.al. (2018) have  
11 brought to the attention of the international meteorology  
12 community a paper relating low latitude basic flow and the  
13 occurrence of typhoons written in Chinese and published already in  
14 1963 that shows MJOs in the zonal wind during the 1958 – 1960  
15 period. In the paper, which is not listed in Table 1 but is now a  
16 part of the history, Xei et.al. (1963) observed that the u-wind  
17 exhibited an oscillatory period of about one and a half month.

18  
19  
20 *Table 1 Chronology of Some Relevant Papers Ordered by*  
21 *Submission Date*

<b>Paper</b>	<b>Date Submitted</b>	<b>Date Published</b>
Cooley and Tukey (1965)	17 August 1964	April 1965
Matsuno (1966)	15 November 1965	February 1966
Julian (1966)	6 December 1965	May 1966
Yanai and Maruyama (1966)	19 July 1966	October 1966
Longuet-Higgins (1968)	28 November 1966	February 1968
Julian (1967)	1 February 1967	September 1967
Maruyama (1967)	28 April 1967	October 1967
Wallace and Kousky (1968)	1 February 1968	September 1968
Yanai et. al. (1968)	26 February 1968	August 1968
Wallace and Chang (1969)	24 February 1969	September 1969
Wallace (1969)	?	October 1969
Wallace (1971)	December 1970	August 1971
Madden and Julian (1971)	21 December 1970	July 1971
Julian (1971)	23 February 1971	December 1971
Madden and Julian (1972)	6 April 1972	September 1972

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41 **4) Epilogue**

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43 **4-1) Developments Related to Two of Our Conclusions**  
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1 The clouds in Figure 7 were based on circumstantial evidence.  
2 During the decade after MJ72, published papers using  
3 wavenumber-frequency analysis of satellite brightness data  
4 (Gruber, 1974; Zangvil, 1975); case studies and spectral analysis  
5 (Yasunari, 1979; 1980); and compositing (Julian and Madden,  
6 1981), provided evidence of cloud behavior consistent with Fig. 7

7  
8 Secondly, we concluded in MJ71 that the spectral results  
9 suggested that the v-wind was not involved in the oscillation.  
10 Specifically, coherence squares between the v-wind and u-wind  
11 were not significantly different from zero. Fifteen years later, we  
12 learned that u and v are coherent and out-of-phase in Northern  
13 Winter and coherent and in-phase in Northern Summer. This in-  
14 and out-of-phase switch between seasons resulted in small  
15 cospectra and a resulting small coherence when, in MJ71, we  
16 averaged over the entire year. The seasonal phase variations are  
17 consistent with surges in the wind from summer to winter  
18 hemispheres (see arguments in Madden, 1986).

#### 19 20 **4-2) Acceptance of MJ71 and MJ72**

21  
22 From 1972 through 1979, MJ71 and MJ72 were cited 17 and 19  
23 times respectively according to the Web of Science. It is interesting  
24 to note that five of the MJ71 citations did not mention the  
25 oscillation itself, but rather they referenced the spectral analysis  
26 method or Julian's discussion about posterior statistical tests.

27  
28 Interest in the two papers picked up in the 1980s when MJ71 and  
29 MJ72 were cited 136 and 140 times respectively. A circumstance  
30 that led to increased interest was the summer MONEX experiment  
31 during May through July of 1979. The MJO was active during that  
32 period. (e.g. Krishnamurti and Subrahmanyam, 1982).

#### 33 34 **5) Conclusions**

35  
36 The path to the initial description of the MJO starts with the  
37 discovery of the QBO in 1961. The QBO stimulated studies aimed  
38 at explaining its remarkable behavior. These studies applied  
39 spectral analysis in innovative ways to describe tropical waves.  
40 The availability of relevant data was increasing along with the  
41 computer power needed for efficient analyses. The fast Fourier  
42 transform which sped up spectral calculations was first coded for  
43 computers in mid-1960. The descriptions contained in MJ71 and  
44 MJ72 relied on the power of spectral analysis.

45



1 Considerable advances in understanding of the MJO have been  
2 made in the 50 intervening years since MJ72. For up to date  
3 information look at:

4 MJO Task Force – a Working Group of the WMO  
5 MJO work at the United States Climate Prediction Center  
6 MJO monitoring and research at Australian BOM

7  
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14 Julian's knowledge, insights, and generosity MJ71 and MJ72 would  
15 not have been written.

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